Carbon pumps in the Benguela upwelling system

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Introduction

- The biological carbon pump influences the flux of carbon dioxide (CO₂) between the ocean and the atmosphere by the photosynthesis of organic carbon (C_{ore}) and the precipitation of calcium carbonate (C_{CaCO3}) . Since these processes have counteracting effects, the CO₂ uptake efficiency of the biological pump depends on the relative importance of these two processes to each other.
- The Benguela upwelling system (BUS) off the southwest African coast is one of the ocean's most productive regions. Contrary to the adjacent South Atlantic where the export of organic carbon is dominated by calcifying plankton^[1], most of the northern BUS (NBUS) shelf is almost free of carbonate deposits^[2] (Fig. 1) and only little is known about particle fluxes on the NBUS shelf.
- To assess whether the CaCO₃-depleted surface sediments on the shelf are caused by low carbonate production in the surface and/or the dissolution of carbonates in the water column we investigated the spatial and temporal variability of the carbonate system (A_p DIC , fCO_2 , $\Omega_{\rm cal}$), the distribution of coccolithophorids, and the annual particle flux on the shelf (HydroBios sediment trap) in 2010.

Fig. 1: Map of NBUS showing the CaCO₃ content of surface
sediment and the location of the sediment trap (cross).

Fig. 4: Relative abundance of coccolithophorids and diatoms vs. a) Ω_{Cal} and b) Si(OH)₄ in 0-20 m water depth<mark>.</mark>

Results

Fig. 2: a) Particle fluxes measured on the continental margin3 and on the shelf (black star and cross in
Fig. 2 b); b) vertical distribution of DIC (color coded) and Ω_{cal} (black line) measured off Walvis Bay.

Water column

- The C_{ore} flux on the shelf is dominated by diatoms in contrast to previous studies conducted along the continental margin where the C_{org} flux is dominated by calcifying plankton^[3] (Fig. 2 a).
- The decomposition of diatomaceous organic matter controls the distribution of O₂ and DIC in the water column and lowers Ω_{Cal} in the ascending subthermocline waters to minimum values of $\Omega_{Cal} = 1.2$ (Fig. 2 b).

Surface

Fig. 3.: Spatial distribution of Ω_{Cal} (color coded) and abundance of coccolithophorids
(circles) observed during a) late austral winter and b) austral summer.

- The Ω_{col} ranges between 1.5 5 in coastal surface water with overall lower values occurring during the upwelling season in austral winter (Fig. 3).
- Coccolithophorids dominate over diatoms at Ω_{Cal} > 2.5 associated with silicate concentrations $Si(OH)_4 < 7 \mu$ mol kg⁻¹ (Fig. 4).
- The results suggest that the low $\Omega_{\rm Cal}$ and high Si(OH)₄ concentrations in coastal surface water favor diatoms rather than coccolithophorids which is reflected in low C_{C_ACO3} fluxes on the shelf and explains why most of the shelf area is almost free of carbonate deposits.

Conclusions

- **The decomposition of organic matter which originates from high, diatom dominated organic carbon fluxes** sustains low Ω_{Cal} in ascending sub**thermocline waters.**
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- **The low** Ω_{col} **in upwelled coastal** 5 b) **surface water reduces the carbonate production relative to the high formation and export of organic** 3GO₃ **carbon on the Namibian shelf.**

bevond slope: shelf: $\begin{bmatrix} 5a \\ 5 \end{bmatrix}$ organic carbon organic carbon + carbonate counter pumi pump surface water $\overline{4}$ ²Cal $\overline{3}$ $\overline{2}$ $40 \frac{1}{5}$ $20\frac{1}{5}$
 $\frac{4}{5}$ 5 c) $\frac{0}{10}$ $=$ TOC surface sediment W. % ँ
≛ 50 $\overline{5}$ \overline{S} Ω ϵ $\frac{1}{136}$ $\frac{1}{13}$ $\frac{1}{14.2}$ de M copen ocean $\cos t \rightarrow$ Fig. 5: Mean a) Ω_{CM} , b) Si(OH)₄ in the surface (0-10 m) and c) TOC and content ₍₂)
ht⁽²⁾ off Walvis Bay. CaCO₃ content of th

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