# Carbon pumps in the Benguela upwelling system

Anita Flohr<sup>[1]</sup>, Tim Rixen<sup>[1,2]</sup>, Niko Lahajnar<sup>[2]</sup>, Anja van der Plas<sup>[3]</sup>, Anja Hansen<sup>[4]</sup>, Norbert Wasmund<sup>[4]</sup>

<sup>[1]</sup> Leibniz Center for Tropical Marine Ecology, Bremen, anita.flohr@zmt-bremen.de; <sup>[2]</sup> Institute for Biogeochemistry and Marine Chemistry, University of Hamburg; <sup>[3]</sup> NatMIRC, Swakopmund, Namibia; <sup>[4]</sup> Leibniz Institute for Baltic Sea Research, Warnemünde

## Introduction

- The biological carbon pump influences the flux of carbon dioxide (CO<sub>3</sub>) between the ocean and the atmosphere by the photosynthesis of organic carbon ( $C_{org}$ ) and the precipitation of calcium carbonate ( $C_{CaCO3}$ ). Since these processes have counteracting effects, the CO<sub>2</sub> uptake efficiency of the biological pump depends on the relative importance of these two processes to each other.
- The Benguela upwelling system (BUS) off the southwest African coast is one of the ocean's most productive regions. Contrary to the adjacent South Atlantic where the export of organic carbon is dominated by calcifying plankton<sup>[1]</sup>, most of the northern BUS (NBUS) shelf is almost free of carbonate deposits<sup>[2]</sup> (Fig. 1) and only little is known about particle fluxes on the NBUS shelf.
- To assess whether the CaCO<sub>3</sub>-depleted surface sediments on the shelf are caused by low carbonate production in the surface and/or the dissolution of carbonates in the water column we investigated the spatial and temporal variability of the carbonate system ( $A_{p}$  DIC , fCO<sub>2</sub>,  $\Omega_{cal}$ ), the distribution of coccolithophorids, and the annual particle flux on the shelf (HydroBios sediment trap) in 2010.





### Results



Fig. 2: a) Particle fluxes measured on the continental margin<sup>[3]</sup> and on the shelf (black star and cross in Fig. 2 b); b) vertical distribution of DIC (color coded) and  $\Omega_{Cal}$  (black line) measured off Walvis Bay.

#### Water column

- The C<sub>org</sub> flux on the shelf is dominated by diatoms in contrast to previous studies conducted along the continental margin where the  $\mathrm{C}_{\mathrm{org}}$  flux is dominated by calcifying plankton<sup>[3]</sup> (Fig. 2 a).
- The decomposition of diatomaceous organic matter controls the distribution of  $O^{}_2$  and DIC in the water column and lowers  $\Omega^{}_{\text{Cal}}$  in the ascending subthermocline waters to minimum values of  $\Omega_{Cal}$  = 1.2 (Fig. 2 b).

#### Fig. 3.: Spatial distribution of $\Omega_{cal}$ (color coded) and abundance of coccolithophorids (circles) observed during a) late austral winter and b) austral summer. Surface

Fig. 4: Relative abundance of coccolithophorids and diatoms vs. a)  $\Omega_{cal}$  and b) Si(OH), in 0-20 m water depth.

- The  $\Omega_{cal}$  ranges between 1.5 5 in coastal surface water with overall lower values occurring during the upwelling season in austral winter (Fig. 3).
- Coccolithophorids dominate over diatoms at  $\Omega_{Cal}$  > 2.5 associated with silicate concentrations Si(OH)<sub>4</sub> < 7  $\mu$ mol kg<sup>-1</sup> (Fig. 4).
- The results suggest that the low  $\Omega_{Cal}$  and high Si(OH)<sub>4</sub> concentrations in coastal surface water favor diatoms rather than coccolithophorids which is reflected in low C<sub>caco3</sub> fluxes on the shelf and explains why most of the shelf area is almost free of carbonate deposits.
- Conclusions The decomposition of organic matter
  - which originates from high, diatom dominated organic carbon fluxes sustains low  $\Omega_{Cal}$  in ascending subthermocline waters.
- The low  $\Omega_{Cal}$  in upwelled coastal surface water reduces the carbonate production relative to the high formation and export of organic carbon on the Namibian shelf.



 The low carbonate formation indicates an efficient upwellingdriven CO<sub>2</sub> uptake by the coastal biological pump.

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